Dissolved organic Fe-binding ligand data from the FRidge (GA13) expedition on RRS James Cook (cruise JC156) from December 2017 to February 2018

Website: <https://www.bco-dmo.org/dataset/923065> **Data Type**: Cruise Results **Version**: 1 **Version Date**: 2024-03-21

Project

» Are strong ligands and dissolved iron tightly coupled in [hydrothermal](https://www.bco-dmo.org/project/917850) systems? (organic iron ligands in hydrothermal systems)

Abstract

Samples were collected as part of the 2017-2018 U.K. GEOTRACES GA13 section cruise along the Mid-Atlantic Ridge (Cruise JC156 on RRS James Cook). Water samples from 11 venting and near venting locations were collected using a Seabird 911 conductivity, temperature, and depth (CTD) titanium rosette using conducting Kevlar wire with an oxidation-reduction potential (ORP) sensor to detect plumes. Teflon coated OTE (Ocean Test Equipment) bottles were pressurized to approximately 7 psi with 0.2 μm filtered air using an oil-free compressor. A Sartobran 300 (Sartorius) filter capsule (0.2 μm) was used to collect filtered seawater samples into clean 250 mL LDPE sample bottles. Bottles and caps were rinsed 3 times with the filtered sample before being filled. Samples were stored frozen at -20°C for organic iron-binding ligand characterization by voltammetry using competitive ligand exchange adsorptive cathodic stripping voltammetry. All dissolved organic iron-binding ligand samples were measured with a BASi controlled growth mercury electrode with an Ag/AgCl¬- reference electrode and platinum auxiliary electrode (Bioanalytical Systems Incorporated) using previously established methods for forward titrations (Buck et al., 2015, 2018; Bundy et al., 2018; Abualhaija and van den Berg, 2014; Hawkes et al., 2013 (Planet. Sci. Lett.)). Reverse titrations (Hawkes et al., 2013 (Anal. Chim. Acta)) were completed on 10 samples from Broken Spur, and TAG hydrothermal vent fields with dissolved iron concentrations between 19.01-90.25 nM.

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Coverage

Location: North Atlantic Mid-Atlantic Ridge (~ 29N, -40E) **Spatial Extent**: **N**:37.29 **E**:-32.28 **S**:26.03 **W**:-45.12

Methods & Sampling

Sampling and cruise transect

Samples were collected as part of the 2017-2018 U.K. GEOTRACES GA13 section cruise along the Mid-Atlantic Ridge. Water samples from 11 venting and near venting locations were collected using a Seabird 911 conductivity, temperature, and depth (CTD) titanium rosette using conducting Kevlar wire with an oxidationreduction potential (ORP) sensor to detect plumes. Teflon coated OTE (Ocean Test Equipment) bottles were pressurized to approximately 7 psi with 0.2 micrometers (μm) filtered air using an oil-free compressor. A Sartobran 300 (Sartorius) filter capsule (0.2 μm) was used to collect filtered seawater samples into clean 250 milliliter (mL) LDPE sample bottles. Bottles and caps were rinsed 3 times with the filtered sample before being filled. Samples were stored frozen at -20 degrees Celsius (°C) for organic iron-binding ligand characterization by voltammetry and mass spectrometry.

Iron-binding ligand concentration and binding strengths measured with competitive ligand exchange-adsorptive cathodic stripping voltammetry

Iron-binding ligand concentrations and binding strengths were determined by competitive ligand exchangeadsorptive cathodic stripping voltammetry (CLE-ACSV) with a BASi controlled growth mercury electrode (CGME) with an Ag/AgCl- reference electrode and platinum auxiliary electrode (Bioanalytical Systems Incorporated). Using previously established methods (Buck et al., 2015, 2018; Bundy et al., 2018; Abualhaija and van den Berg, 2014; Hawkes et al., 2013 (Planet. Sci. Lett.)). 40 frozen filtrate (<0.2 µm) samples with dissolved iron concentrations between 0.41-11.67 nanomolar (nM) (Table S1-S2 of Hoffman et al., 2023) were thawed in a 4°C fridge prior to analysis. A 15-point titration curve was analyzed for each sample. Briefly, within each titration, every point sequentially received 10 mL of sample, 7.5 micromolar (µM) of borate-ammonium buffer, 10 µM salicylaldoxime (SA) as the added ligand, and a dissolved Fe addition. Data was collected using the Epsilon Eclipse Electrochemical Analyzer (v.213) with a deposition time of 120 seconds and analyzed using ElectroChemical Data Software (v2001-2014) and ProMCC (v2008-2018) to determine peak areas and ironbinding ligand parameters, respectively.

Reverse Titration-CLE-ACSV

Reverse titration-CLE-ACSV (RT-CLE-ACSV) (Hawkes et al., 2013 (Anal. Chim. Acta)) was completed on 10 samples from Broken Spur, and TAG hydrothermal vent fields with dissolved Fe concentrations between 19.01- 90.25 nM (Table S1-S2 of Hoffman et al., 2023). Briefly, a 10-point titration curve was analyzed for each sample with each titration point consisting of 10 mL of sample buffered with 7.5 µM boric acid and the competitive ligand 1-nitroso-2-napthol (NN). All samples were analyzed on a BASi Controlled Growth Mercury Electrode (CGME) with the Epsilon Eclipse Electrochemical Analyzer (v.213) and deposition time of 120 seconds. For each sample, the competitive ligand NN was added in concentrations of 0.5, 1, 2, 3, 4, 6, 9, 15, 20, and 40 µM. Samples were equilibrated overnight and purged with N₂ (99.99% UHP) for 5 minutes before analysis. At the end of each titration, three dissolved iron additions (3-15 nM) were added to the final titration point to get the total concentration of iron in equilibrium with the natural ligands. Data was analyzed using ElectroChemical Data Software (v2001-2014) to acquire peak areas and a package in R using the model parameters of $\beta_{FennN3} = 5.12 \times 10^{16}$, Xmin = 0.8, Xmax = 0.9, and c1high = 0.75 to determine the iron-binding ligand parameters (Hawkes et al., 2013 (Anal. Chim. Acta)). These parameters were chosen based on the recommendations for undersaturated samples and titrations curves where ipmax was not reached (Hawkes et al., 2013 (Anal. Chim. Acta)). All other parameters within the model we kept at the default values.

Data Processing Description

Forward titrations were analyzed using ElectroChemical Data Software (v2001-2014) and ProMCC (v2008- 2018) to determine peak areas and Fe-binding ligand parameters, respectively.

Reverse titration data was analyzed using ElectroChemical Data Software (v2001-2014) to acquire peak areas and a package in R using the model parameters of $\beta_{FeNN3} = 5.12 \times 10^{16}$, Xmin = 0.8, Xmax = 0.9, and c1high = 0.75 to determine the Fe-binding ligand parameters (Hawkes et al., 2013 (Anal. Chim. Acta)).

Additional data for siderophore analyses have been deposited on Massive under the DOI [10.25345/C5V97ZW7N](http://dx.doi.org/10.25345/C5V97ZW7N).

Microbial 16S rRNA data have been deposited on GenBank under the accession number BioProject #PRJNA865382.

BCO-DMO Processing Description

- Imported original file "FRidge_GA13_CSV_data_final_bcodmo.xlsx" into the BCO-DMO system.
- Flagged "NaN" as a missing data identifier. Missing data are empty/blank in the final CSV file.
- Renamed fields to comply with BCO-DMO naming conventions.
- Saved the final file as "923065 v1 ga13 diss org fe-binding ligands.csv".

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Data Files

File

923065_v1_ga13_diss_org_fe-binding_ligands.csv(Comma Separated Values (.csv), 4.15 KB) MD5:da194bf71e6b69636cd9adff12eac119

Primary data file for dataset ID 923065, version 1

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Related Publications

Abualhaija, M. M., & van den Berg, C. M. G. (2014). Chemical speciation of iron in seawater using catalytic cathodic stripping voltammetry with ligand competition against salicylaldoxime. Marine Chemistry, 164, 60–74. doi[:10.1016/j.marchem.2014.06.005](https://doi.org/10.1016/j.marchem.2014.06.005) **Methods**

Buck, K. N., Sedwick, P. N., Sohst, B., & Carlson, C. A. (2018). Organic complexation of iron in the eastern tropical South Pacific: Results from US GEOTRACES Eastern Pacific Zonal Transect (GEOTRACES cruise GP16). Marine Chemistry, 201, 229–241. https://doi.org[/10.1016/j.marchem.2017.11.007](https://doi.org/10.1016/j.marchem.2017.11.007) Methods

Buck, K. N., Sohst, B., & Sedwick, P. N. (2015). The organic complexation of dissolved iron along the U.S. GEOTRACES (GA03) North Atlantic Section. Deep Sea Research Part II: Topical Studies in Oceanography, 116, 152–165. doi[:10.1016/j.dsr2.2014.11.016](https://doi.org/10.1016/j.dsr2.2014.11.016) Methods

Bundy, R. M., Boiteau, R. M., McLean, C., Turk-Kubo, K. A., McIlvin, M. R., Saito, M. A., Van Mooy, B. A. S., & Repeta, D. J. (2018). Distinct Siderophores Contribute to Iron Cycling in the Mesopelagic at Station ALOHA. Frontiers in Marine Science, 5. https://doi.org[/10.3389/fmars.2018.00061](https://doi.org/10.3389/fmars.2018.00061) **Methods**

Hawkes, J. A., Connelly, D. P., Gledhill, M., & Achterberg, E. P. (2013). The stabilisation and transportation of dissolved iron from high temperature hydrothermal vent systems. Earth and Planetary Science Letters, 375, 280–290. https://doi.org[/10.1016/j.epsl.2013.05.047](https://doi.org/10.1016/j.epsl.2013.05.047) **Methods**

Hawkes, J. A., Gledhill, M., Connelly, D. P., & Achterberg, E. P. (2013). Characterisation of iron binding ligands in seawater by reverse titration. Analytica Chimica Acta, 766, 53–60. https://doi.org[/10.1016/j.aca.2012.12.048](https://doi.org/10.1016/j.aca.2012.12.048) Methods

Hoffman, C. L., Monreal, P. J., Albers, J. B., Lough, A. J. M., Santoro, A. E., Mellett, T., Buck, K. N., Tagliabue, A., Lohan, M. C., Resing, J. A., & Bundy, R. M. (2023). Microbial strong organic ligand production is tightly coupled to iron in hydrothermal plumes. https://doi.org[/10.1101/2023.01.05.522639](https://doi.org/10.1101/2023.01.05.522639) Results

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Related Datasets

IsRelatedTo

Randelle M. Bundy. (2022). MassIVE MSV000090721 - Hydrothermal Siderophore Analysis on UK GEOTRACES GA13 (FRidge) [Data set]. MassIVE. https://doi.org[/10.25345/C5V97ZW7N](https://doi.org/10.25345/C5V97ZW7N)

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Deployments

JC156 Website <https://www.bco-dmo.org/deployment/923071> **Platform RRS** James Cook **Start Date** 2017-12-20 **End Date** 2018-02-01 **Description** See more about this cruise at: https://www.bodc.ac.uk/resources/inventories/cruise_inventory/report/16349/

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Project Information

Are strong ligands and dissolved iron tightly coupled in hydrothermal systems? (organic iron ligands in hydrothermal systems)

NSF Award Abstract:

This award is funded in whole or in part under the American Rescue Plan Act of 2021 (Public Law 117-2).

Iron is one of the most abundant elements in the Earth's crust, but it is extremely diluted in the ocean. Ironpoor surface waters limit the growth of microscopic marine life, called phytoplankton, and their ability to remove carbon from the atmosphere and surface ocean. However, over the last few decades, our understanding of how iron enters the ocean has evolved. Recent data has shown that deep-sea hot springs, also known as hydrothermal vents, impact global iron budgets and are important for surface iron supply. Hydrothermal vents are found globally along volcanic spreading centers where new seafloor is created through tectonic activity. The new porous seafloor allows seawater to circulate through the hot, chemically reactive rocks to create hydrothermal fluids. These fluids are less dense (hotter, 300-400°C) than deep ocean waters (2°C), so the water exiting the vents rises while mixing with ambient seawater, eventually forming hydrothermal plumes. These nutrient-rich plumes can extend for 10-1000s of kilometers into the ocean interior. To account for the long-range transport of hydrothermal iron into the ocean interior, models have shown that stabilizing agents (i.e. organic ligands) are needed to prevent iron from precipitating and settling to the seafloor. However, we still do not know the sources and identities of these organic ligands, as well as how common they are in various hydrothermal systems across the global ocean. Investigating these mechanism(s) for hydrothermal iron stabilization across different vent systems will provide insight into both local and longrange iron utilization by deep-sea marine microorganisms and phytoplankton in the surface ocean.

In this project, the sources, concentration, and identities of iron-binding organic ligands in hydrothermal plumes from four different volcanic spreading centers will be examined to understand their impact on iron stabilization and transport into the ocean interior. The major aim of this research is to test whether (1) the concentrations of strong organic ligands tightly control the distal transport of hydrothermally derived dissolved iron in neutrally buoyant plumes across a variety of hydrothermal vent systems and (2) investigate if microbes from hydrothermal systems are responsible for production of these strong organic ligands (i.e. siderophores). This work will use a combination of existing samples and samples of opportunity that will be collected during an upcoming field expedition, each from distinct spreading centers. These findings would significantly enhance our understanding of hydrothermal iron transport and aid in future modeling efforts on the fate of hydrothermal iron in the global iron cycle. This project will support the training of two early career scientists, an undergraduate intern, and STEM workshop kits for middle school programs about deep-sea environments, which will be developed in collaboration and made freely available through the NOAA Pacific Marine Environmental Education and Outreach webpage.

This award reflects NSF's statutory mission and has been deemed worthy of support through evaluation using the Foundation's intellectual merit and broader impacts review criteria.

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